

# Chaotic advection in point vortex models and two-dimensional turbulence

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(Received 18 May 1993; accepted 17 March 1994)

Physics of Fluids 6, 2465 (1994)



# 20 years of Lagrangian statistics

## Outline

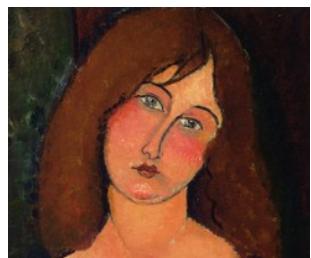


Historical part:

- Lagrangian and Eulerian statistics
- Exit time statistics
- Applications in turbulence

Modern part:

- Active Lagrangian tracers
- Swimming in turbulence



# Prehistory



## Chaotic advection in point vortex models and two-dimensional turbulence

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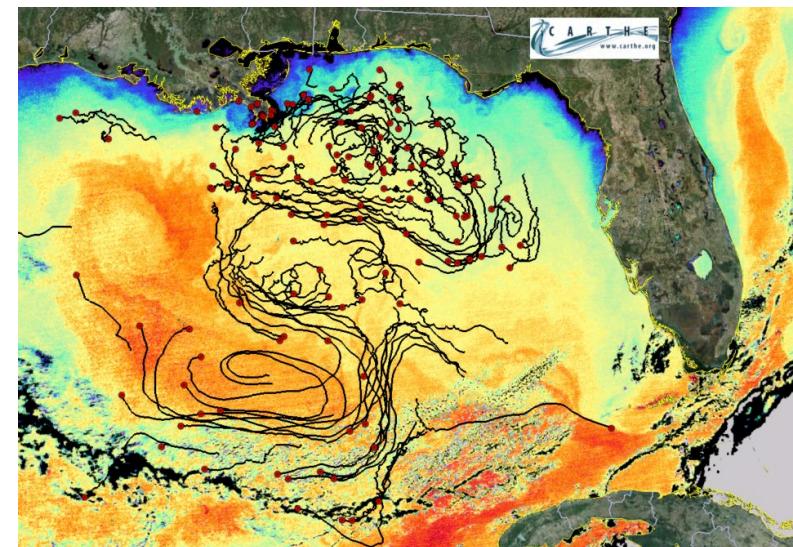
Physics of Fluids 6, 2465 (1994)

General question: how to connect Eulerian and Lagrangian statistics in flows

Relevant for applications: e.g. reconstruct circulation from drifter trajectories

In 2D: a nonautonomous Hamiltonian system

$$\begin{aligned}\frac{dx}{dt} &= +\frac{\partial \psi}{\partial y} & \psi(x, y, t) \text{ is the stream function} \\ \frac{dy}{dt} &= -\frac{\partial \psi}{\partial x}\end{aligned}$$



chaotic trajectories are possible for regular (time periodic) 2d flows (1+1/2 dof)

Regular (non chaotic) trajectories are possible even for chaotic flows

**Point vortex model:** discrete N-particle model for 2D ideal fluid

[Kirchhoff, Onsager...]

$$\omega(\mathbf{x}, t) = \sum_{\alpha=1}^N \Gamma_\alpha \delta(\mathbf{x} - \mathbf{x}_\alpha(t))$$

the motion of point vortices  $\mathbf{x}_\alpha(t)$  is ruled by the Hamiltonian

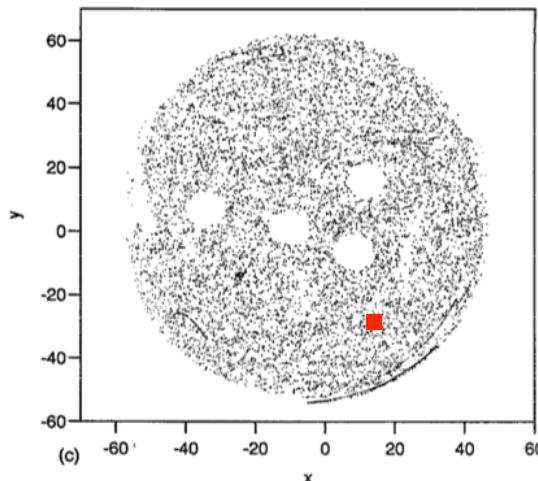
$$H = -\frac{1}{4\pi} \sum_{\alpha \neq \beta} \Gamma_\alpha \Gamma_\beta \ln r_{\alpha\beta}$$



Lars Onsager (1903-1976)

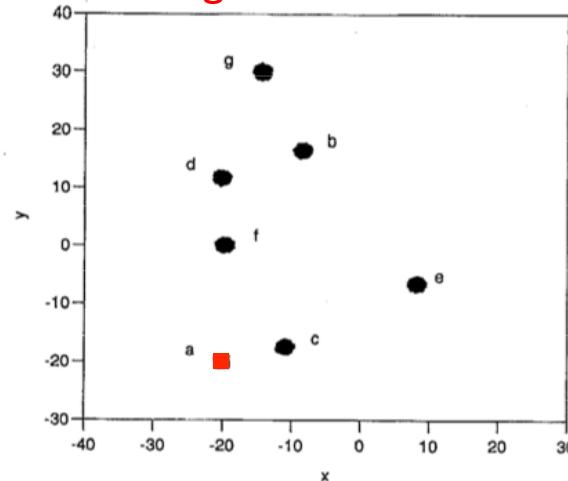
For  $N \geq 4$  the motion of vortices (the **flow**) is chaotic: how is the (**Lagrangian**) motion of a drifter?

chaotic drifters



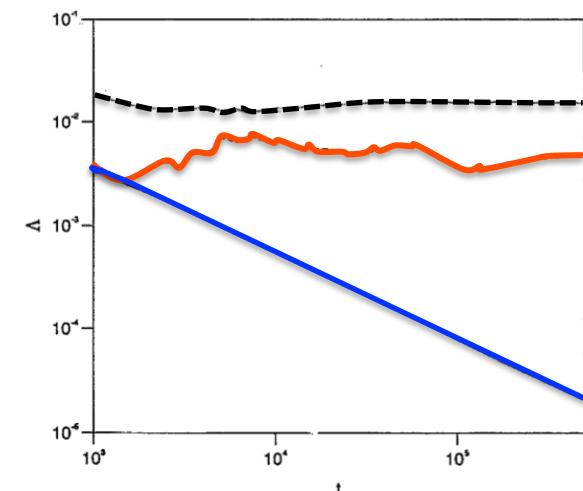
trajectories starting far from vortices visit all the available phase space

regular drifters



trajectories starting close to a vortex remain trapped

Lyapunov exponent



chaotic  $\lambda > 0$  Eulerian:  $\lambda > 0$   
regular  $\lambda = 0$

# Exit time statistics

The Lyapunov exponent measures the rate of separation of **infinitesimally close** trajectories  
In many applications the **growth of finite separations** is more relevant

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## Growth of Noninfinitesimal Perturbations in Turbulence

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(Received 15 April 1996)

Exit time  $T_\rho(R)$ : time for separation to grow from  $R$  to  $\rho R$

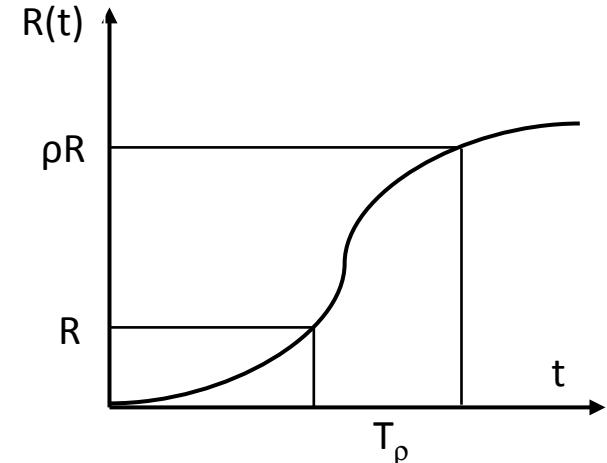
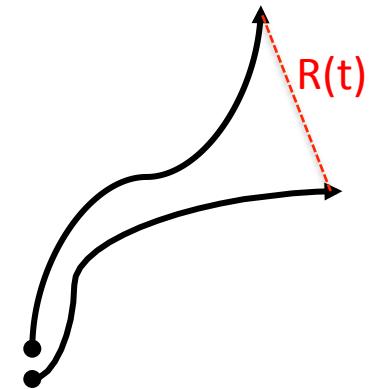
$$R(t + T_\rho) = \rho R(t)$$

Finite size Lyapunov exponent

$$\lambda(R) \equiv \frac{1}{\langle T_\rho(R) \rangle} \ln \rho$$

for infinitesimal separations

$$\lim_{R \rightarrow 0} \lambda(R) = \lambda$$



# Dispersion of passive tracers in closed basins: Beyond the diffusion coefficient

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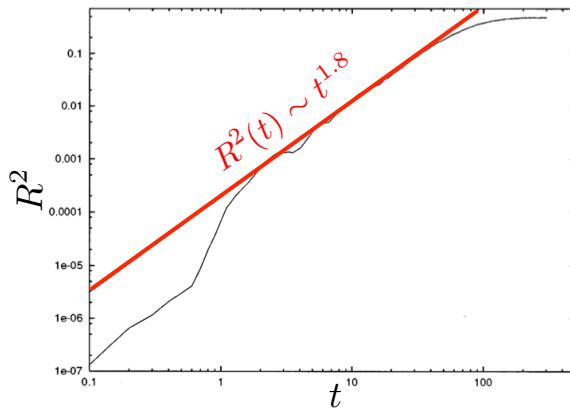
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and Istituto Nazionale Fisica della Materia, Unità di Roma, Piazzale Aldo Moro 5, 00185 Roma, Italy

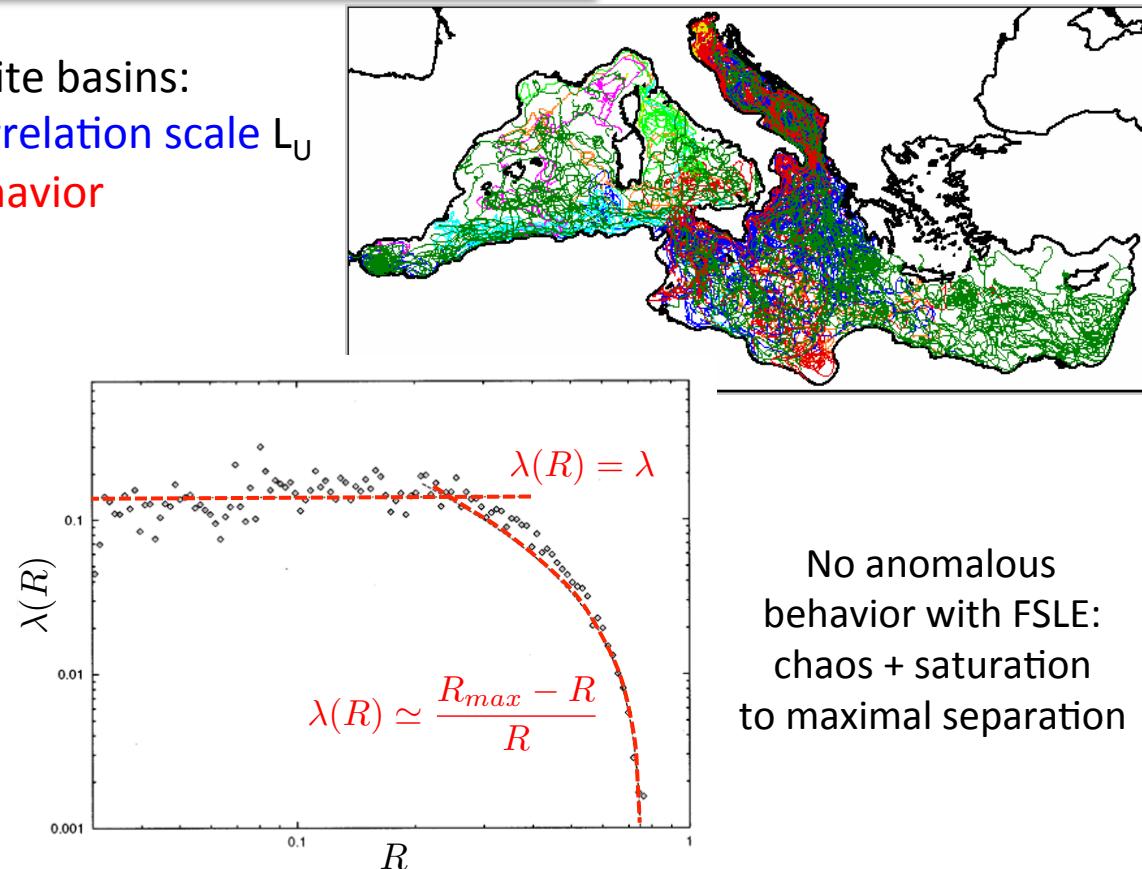


Physics of Fluids 9, 3162(1997)

Dispersion of Lagrangian drifter in finite basins:  
**no scale separation between flow correlation scale  $L_U$**   
**and domain scale  $L_B$  : no diffusive behavior**



Apparent anomalous behavior  
 $\langle R^2(t) \rangle \sim t^{2\nu}$  at intermediate times

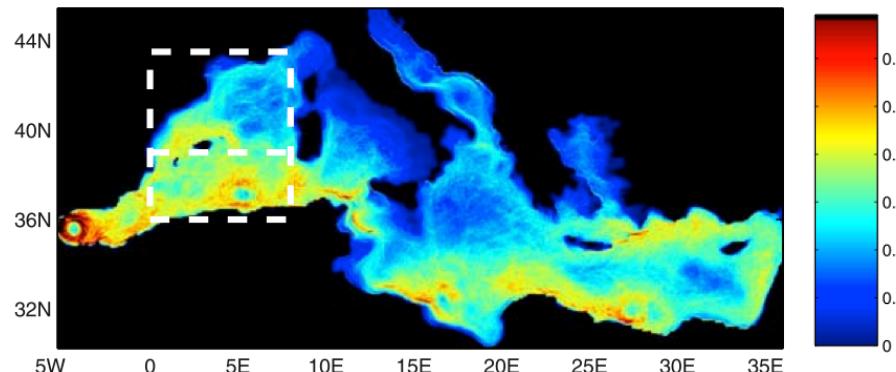


No anomalous behavior with FSLE:  
chaos + saturation  
to maximal separation

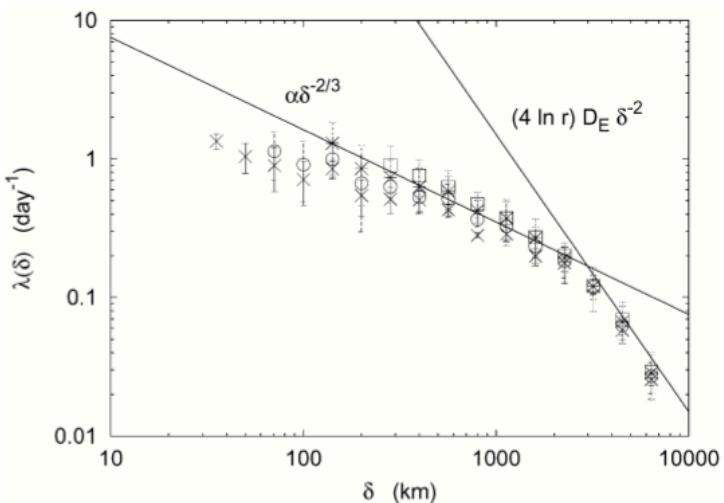
# Use of the Finite Size Lyapunov Exponent

## Mixing properties of the Mediterranean Sea

F. d'Ovidio *et al*, "Mixing structures in the Mediterranean Sea from finite-size Lyapunov exponents"  
Geophys. Res. Lett. **31**, L17203 (2004)



## Relative dispersion and inverse cascade in the stratosphere



## EOLE project: 483 balloons in the southern hemisphere

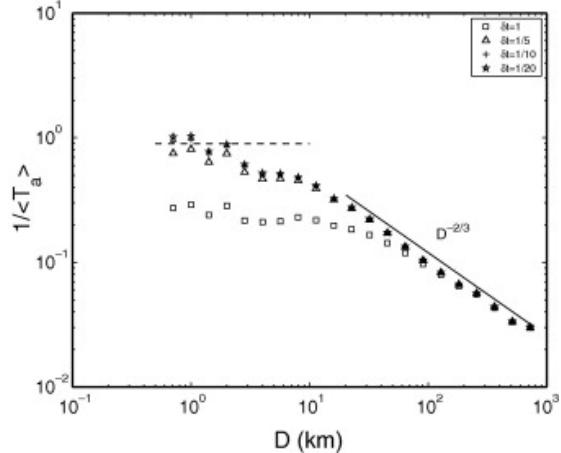
G. Lacorata, E. Aurell, B. Legras, A. Vulpiani,  
"Evidence for a  $k^{5/3}$  Spectrum from the EOLE Lagrangian  
Balloons in the Low Stratosphere", J. Atmos. Sci. 61, 2936 (2004)

# Use of the Finite Size Lyapunov Exponent / 2

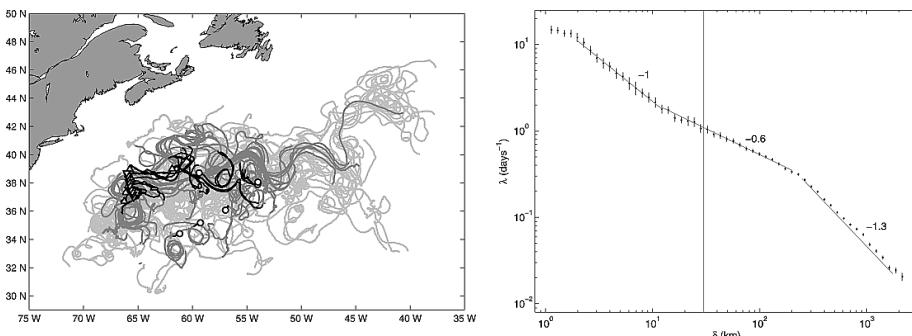
Relative dispersion in the Gulf of Mexico

700 surface drifters deployed at pairs

J.H. LaCasce, "Statistics from Lagrangian observations"  
Prog. Oceanography 77, 1 (2008)



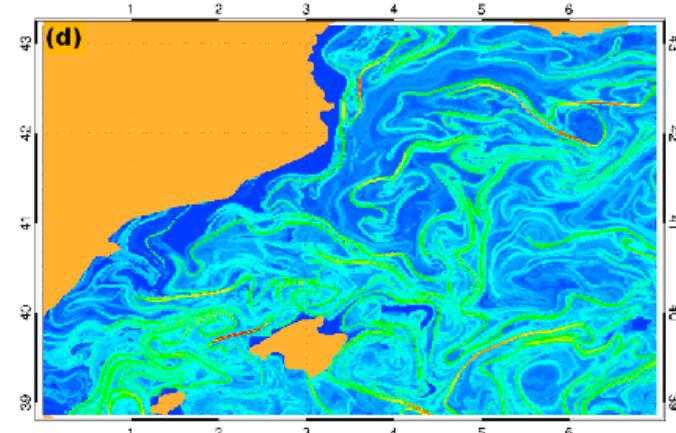
Dispersion of surface drifters in the Gulf stream



R. Lumpkin, S. Eliot, "Surface drifter pair spreading in the North Atlantic", J. Geophys. Res. 115, C12017 (2010)

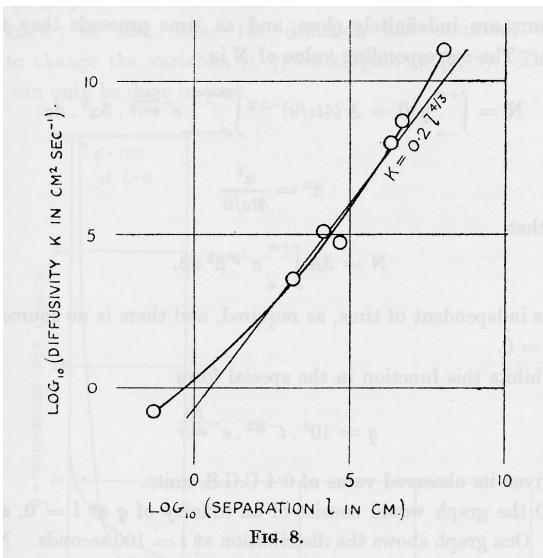
Numerical study of dispersion in the Balearic sea

I. Hernández-Carrasco, C. López, E. Hernández-García, A. Turiel,  
"How reliable are finite-size Lyapunov exponents for the assessment  
of ocean dynamics?" Ocean Mod. 36, 208 (2011)



# Relative dispersion in turbulence

First empirical evidence of Kolmogorov scaling in turbulence (in 1926)



Diffusivity in the atmosphere

$$K(R) \simeq R^{4/3}$$

From Kolmogorov scaling:

$$K(R) = \frac{dR^2}{dt} \simeq R\delta v(R) \simeq \varepsilon^{1/3} R^{4/3}$$



L.F. Richardson (1881-1953)

Diffusion equation for *distance neighbor function*  $p(\mathbf{R}, t)$

$$\frac{\partial p(\mathbf{R}, t)}{\partial t} = \frac{\partial}{\partial R_i} \left[ K(R) \frac{\partial p(\mathbf{R}, t)}{\partial R_i} \right]$$

Solution with  $\delta$  initial condition:  $p(\mathbf{R}, t) = \frac{C}{(\alpha t)^{9/2}} \exp \left( -\frac{9R^{2/3}}{4\alpha t} \right)$

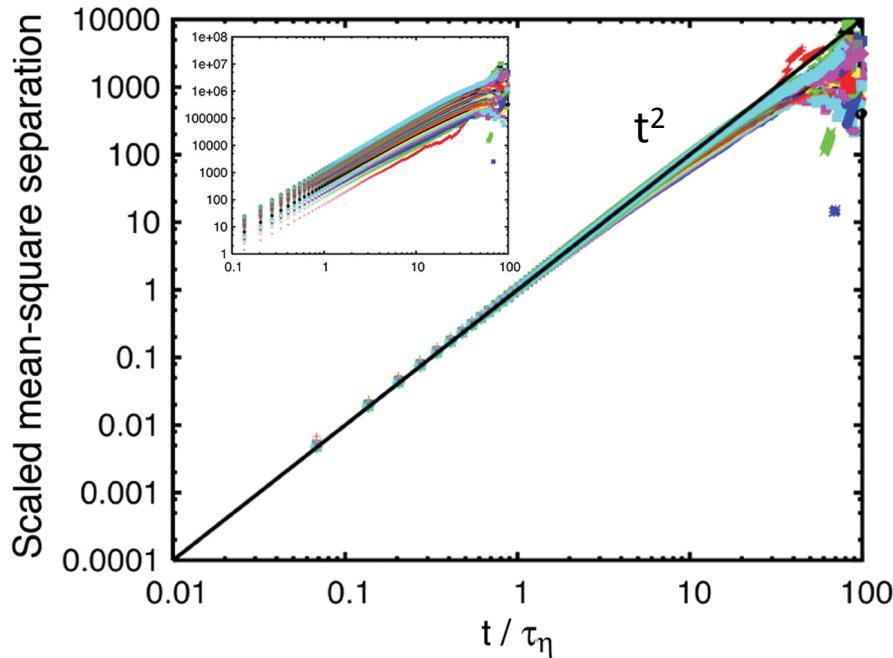
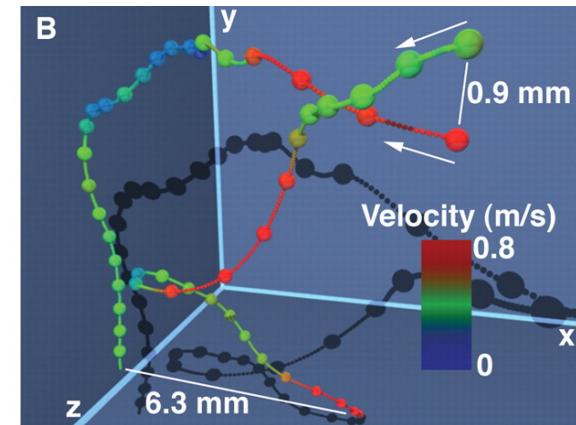
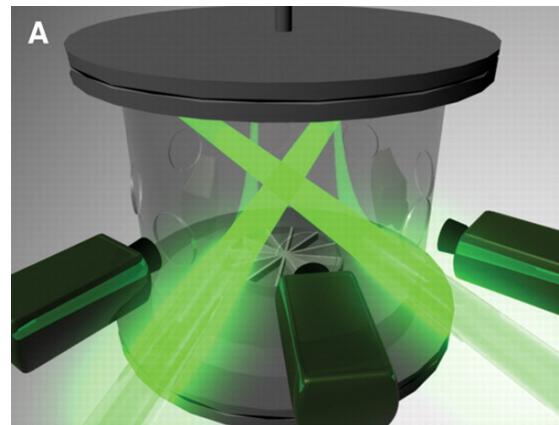
**Explosive** separation of trajectories:

$$\langle R^2(t) \rangle = g\varepsilon t^3$$

(Richardson's law)

# Exit times in turbulence

Difficulties in observing Richardson's law because of strong fluctuations at small scales



Recent laboratory experiments at high Reynolds numbers failed to observe Richardson's  $t^3$  law (they observe  $t^2$ , ballistic)

# Exit times in turbulence

Difficulties in observing Richardson's law because of strong fluctuations at small scales

## Pair dispersion in synthetic fully developed turbulence

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A. Crisanti and A. Vulpiani

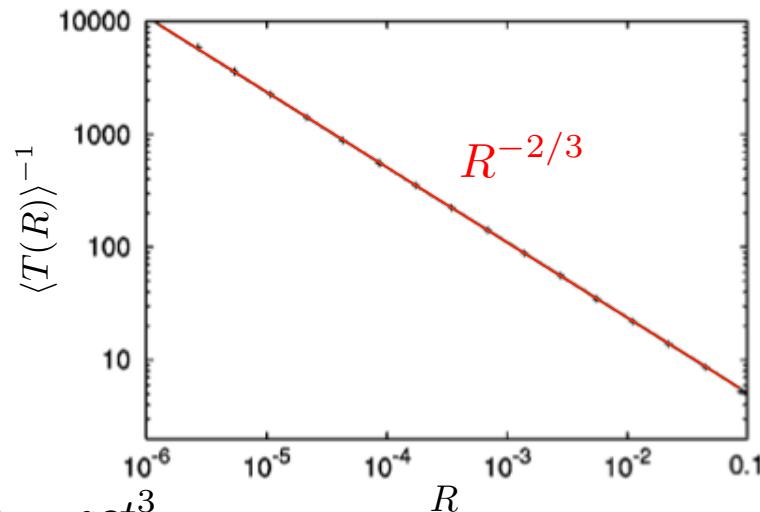
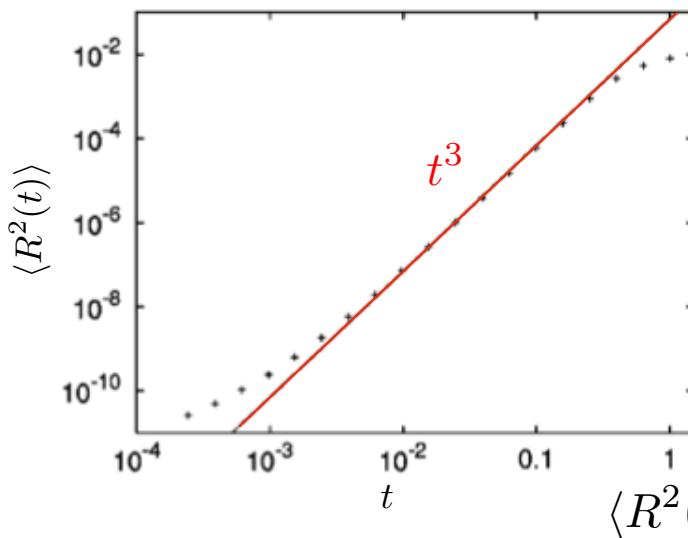
Dipartimento di Fisica, Università di Roma "La Sapienza," Piazza le Aldo Moro 2, 00185 Roma, Italy  
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(Received 1 June 1999)

Phys. Rev. E 60,  
6734 (1999)



Use the exit time statistics to avoid contamination between different scales



Synthetic turbulence

much  
better  
scaling

# Exit time statistics in direct numerical simulations of Navier-Stokes equations

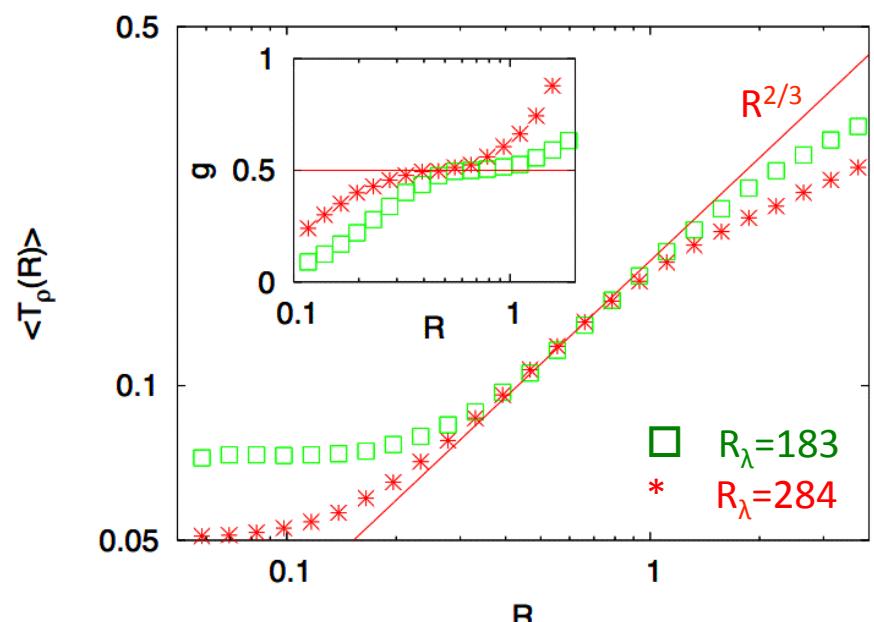
Assuming the Richardson model for relative dispersion  
the mean exit time is given by a first passage problem  
as

$$\langle T_\rho(R) \rangle = \frac{3}{2d} \frac{\rho^{2/3} - 1}{\alpha \varepsilon^{1/3} \rho^{2/3}} R^{2/3}$$

Richardson constant

$$g = \frac{143}{81} \frac{(\rho^{2/3} - 1)^3}{\rho^2} \frac{R^2}{\varepsilon \langle T_\rho(R) \rangle^3} \simeq 0.5$$

$$\langle R^2(t) \rangle = g \varepsilon t^3$$



# 20 years of Lagrangian statistics

## Outline



Modern part:

- Active Lagrangian tracers
- Swimming in turbulence

In collaboration with

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R. Stocker, M. Barry (MIT, Cambridge, USA)

W.M. Durham (University of Oxford, UK), E. Climent (CNRS Toulouse, FR)

W.M. Durham et al, *Nature Comm* **4**, 2148 (2013)

F. De Lillo et al, *Phys Rev Lett* **112**, 044502 (2014).

# Phytoplankton: Active Lagrangian Tracers

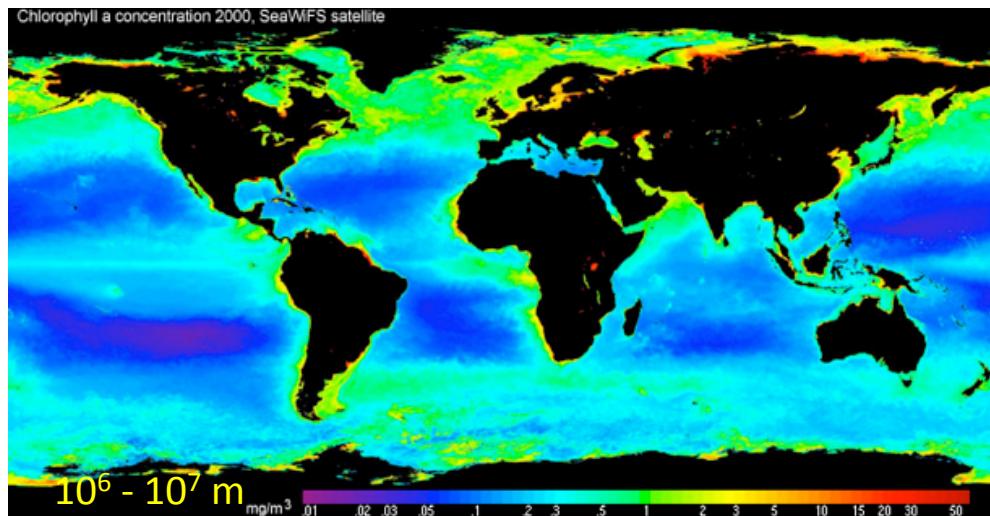
- unicellular organisms of many forms and sizes between 1 and 100 µm
- about 5000 species
- about 50% of photosynthetic activity on Earth
- at the basis of the marine food web
- harmful algal blooms from toxic species
- many species are able to swim
- patchiness at different scales

Turbulence mediates many processes crucial to the ecology of phytoplankton, including motility, nutrient uptake, and cell-cell encounters.

# Motivations

Collective phenomena in micro-swimmers

- plankton patchiness over many scales
- (toxic) algal blooms and “thin layers”

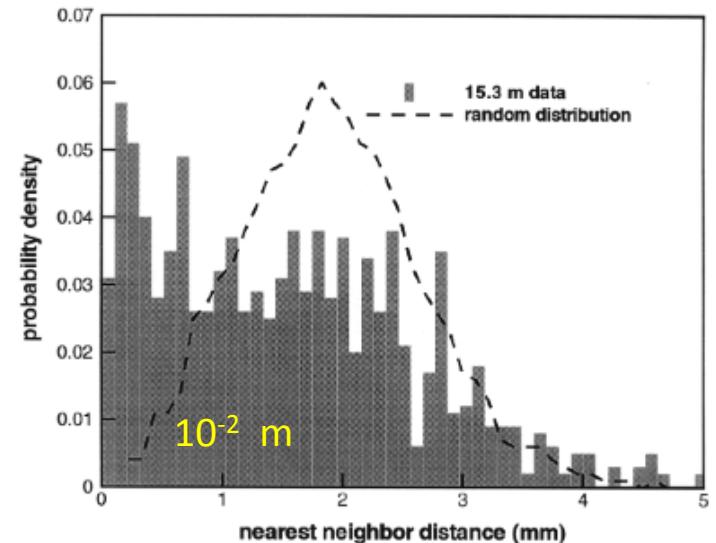


Motile phytoplankton is found to be **more patchy** than non-motile one

E.Malkiel, O.Alquaddoomi, J.Katz  
Meas. Sci. Technol. **10** (1999)



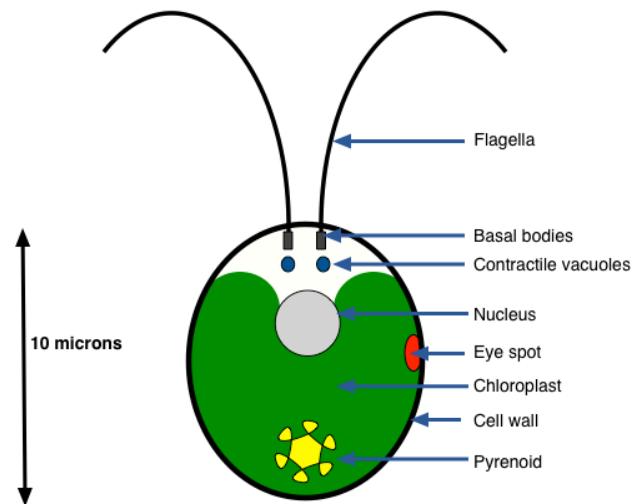
Red tide of *Karenia brevis*



**Figure 10.** Measured nearest-neighbour distances and expected values from a random distribution.

# Swimming algae

The genus *Chlamydomonas*, unicellular flagellate, a model organism for molecular biology.

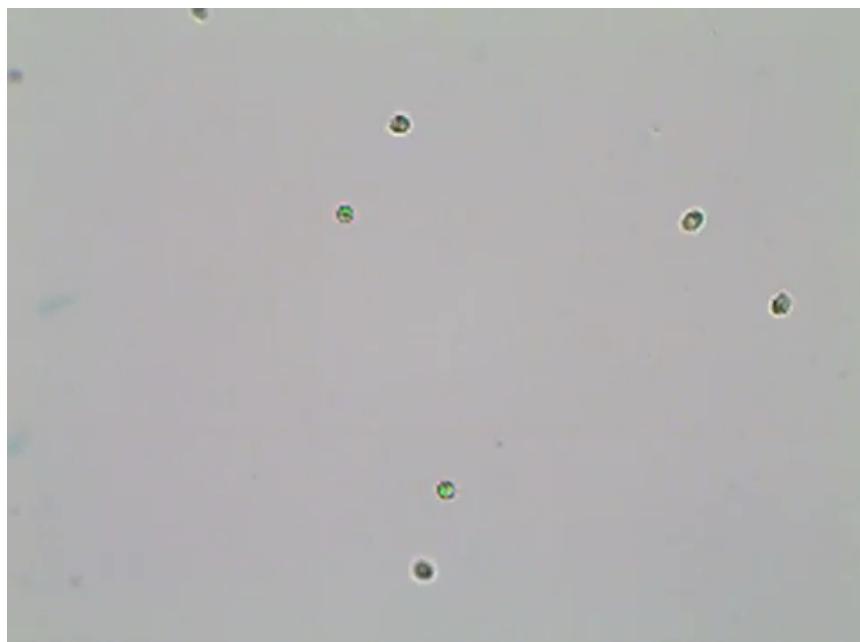
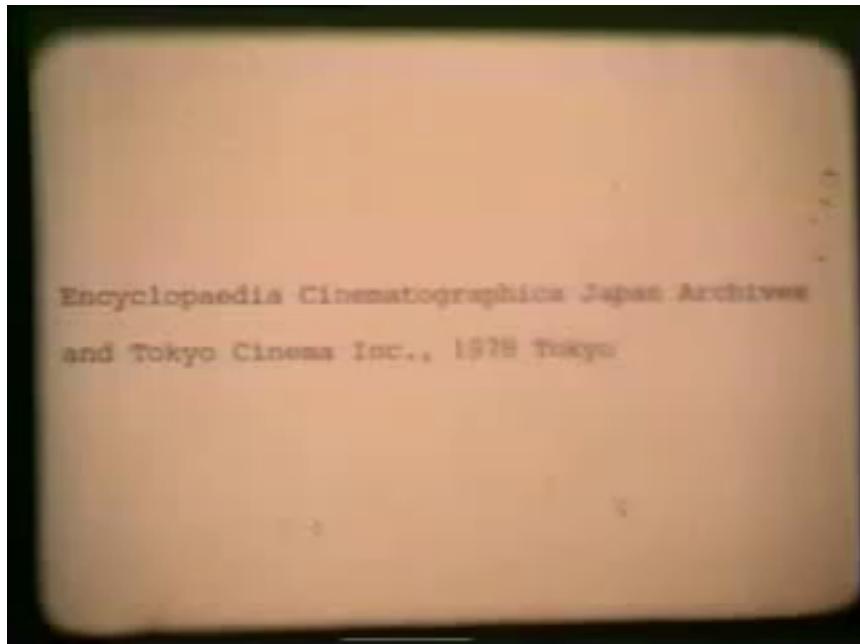


Both sexual and asexual reproduction

Eyespot: light receptor for phototactic (and photophobic) behaviors  
(cell rotates at about 2 Hz to detect direction)

Slightly heavy, **unbalanced weight** (bottom heavy due to chloroplast distribution): they naturally **swim against gravity (gyrotaxis)**

**Good swimmers:** about 10 body-length per second



# Gyrotactic model for bottom-heavy cells

$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + v_s \mathbf{p}$$

$$\frac{d\mathbf{p}}{dt} = -\frac{1}{2v_o} [\mathbf{A} - (\mathbf{A} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$

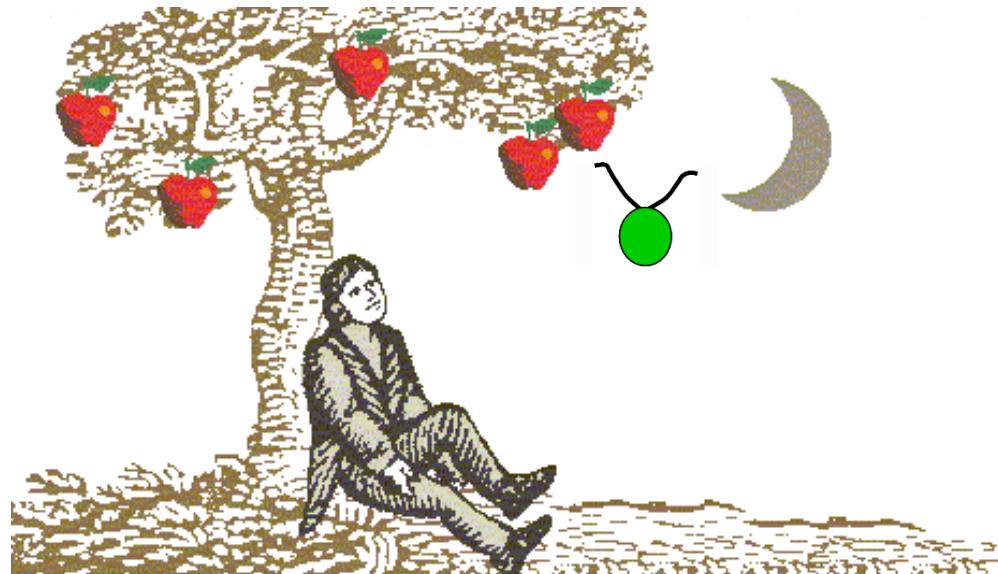
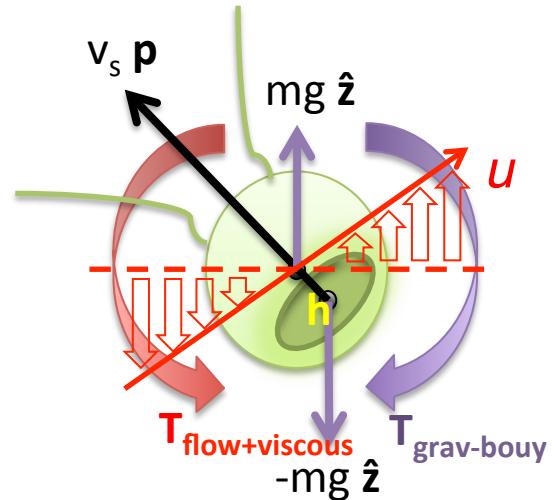
$$\mathbf{A} = \mathbf{g} - \mathbf{a}(\mathbf{x}, t)$$

$$v_s \simeq 100 \mu\text{m s}^{-1}$$

$$v_0 = \frac{3\nu}{h} = O(1)m s^{-1}$$

JO Kessler, Nature **313**, 218 (1985)

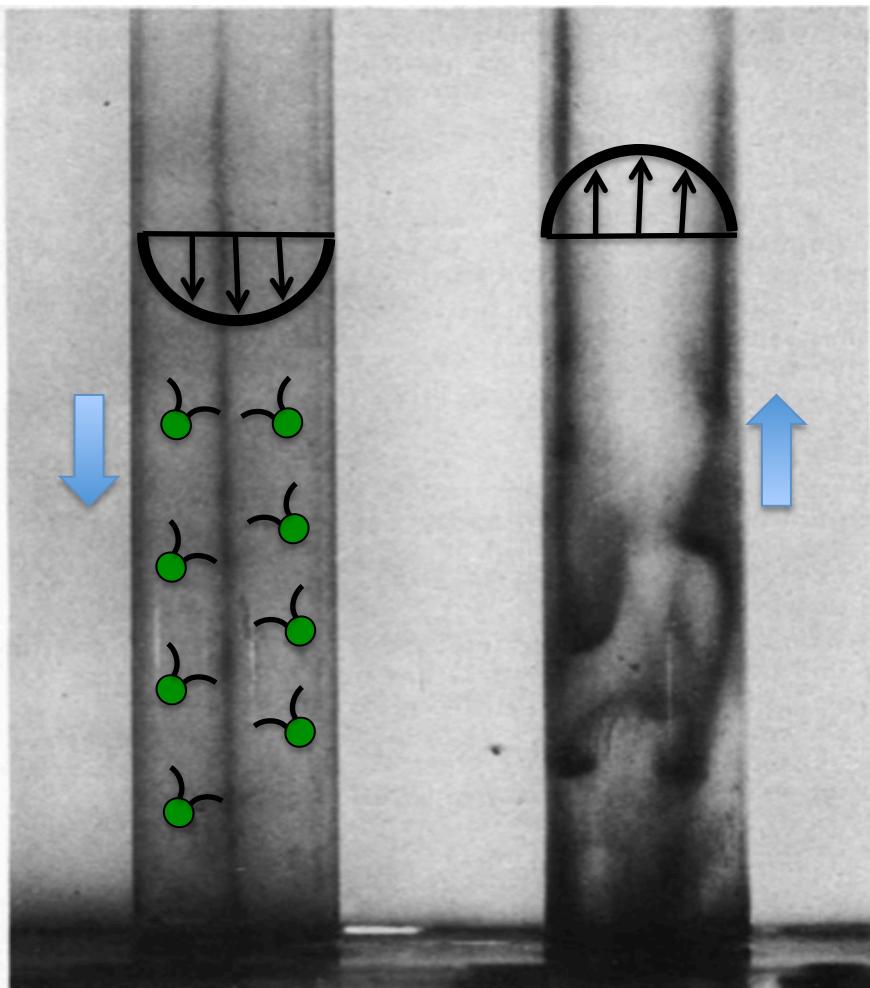
T.J. Pedley, J.O. Kessler, Proc. Roy Soc. B **231** (1987)



# Gyrotactic focusing in (laminar) pipe flows

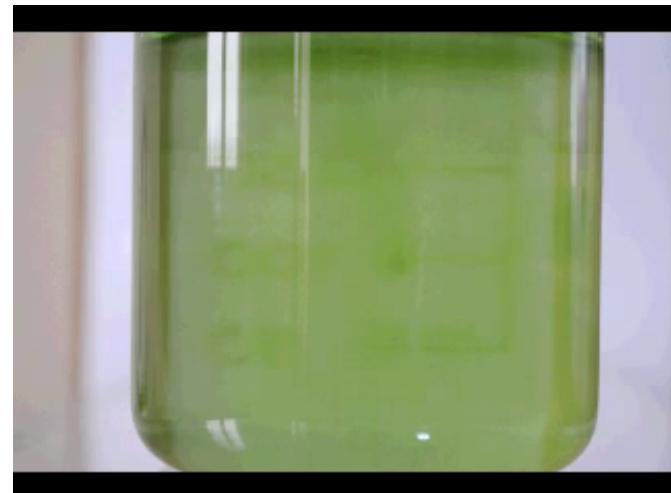
$$a = 0$$

$$\frac{d\mathbf{p}}{dt} = -\frac{1}{2v_o} [\mathbf{g} - (\mathbf{g} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$



Swimming cells concentrate  
in the center of downwelling flows

Concentration in bioconvection



JO Kessler, Nature 313, 218 (1985)

# Acceleration in turbulence

$$\langle a^2 \rangle = a_0 \epsilon^{3/2} \nu^{-1/2}$$

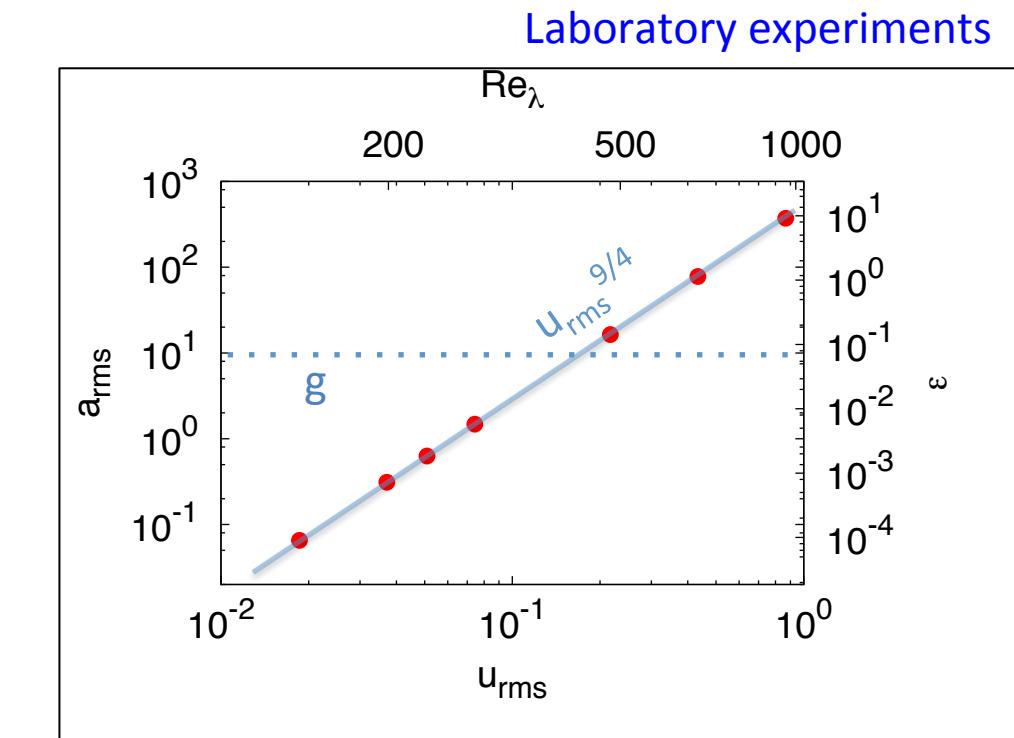
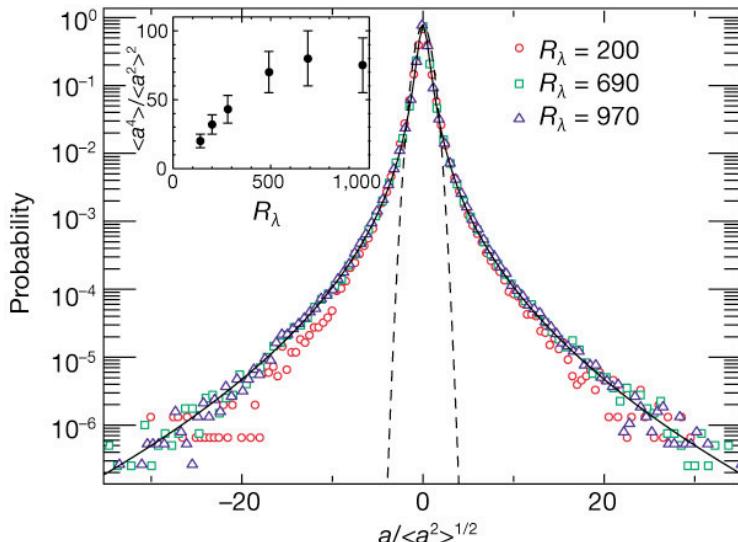
with  $a_0 \simeq 5 - 6$  and  $\nu_{H_2O} \simeq 10^{-6} m^2 s^{-1}$

Kinetic energy dissipation in the oceanic mixing layer  $\epsilon \leq 10^{-4} m^2 s^{-3}$

which means

$$a_{rms} \simeq 0.1 ms^{-2} \ll g$$

Experimental and numerical data show that local acceleration in turbulence is **extremely intermittent** with PDF with very wide tails



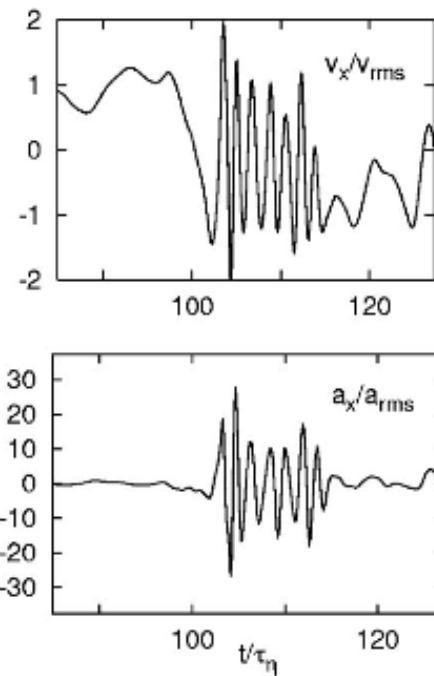
from Voth et al, JFM 469, 121 (2002)

Even if  $a_{rms} < g$ , local fluid acceleration can exceed gravity

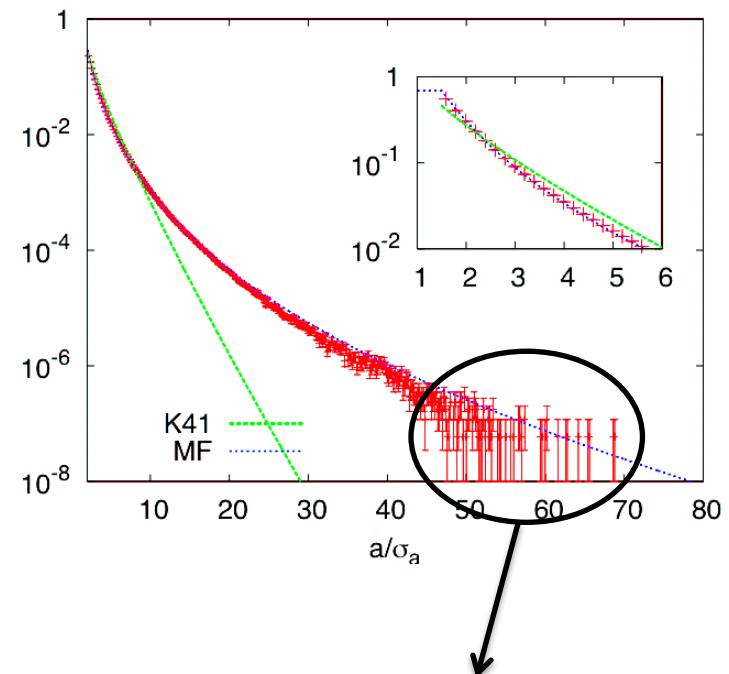
La Porta et al, Nature 409, 1017 (2001)

# Origin of extreme accelerations: turbulent vortices

Trapping of particles in small scale vortices



Maximal acceleration inside vortices



# The effects of fluid accelerations: a toy model for a vortex

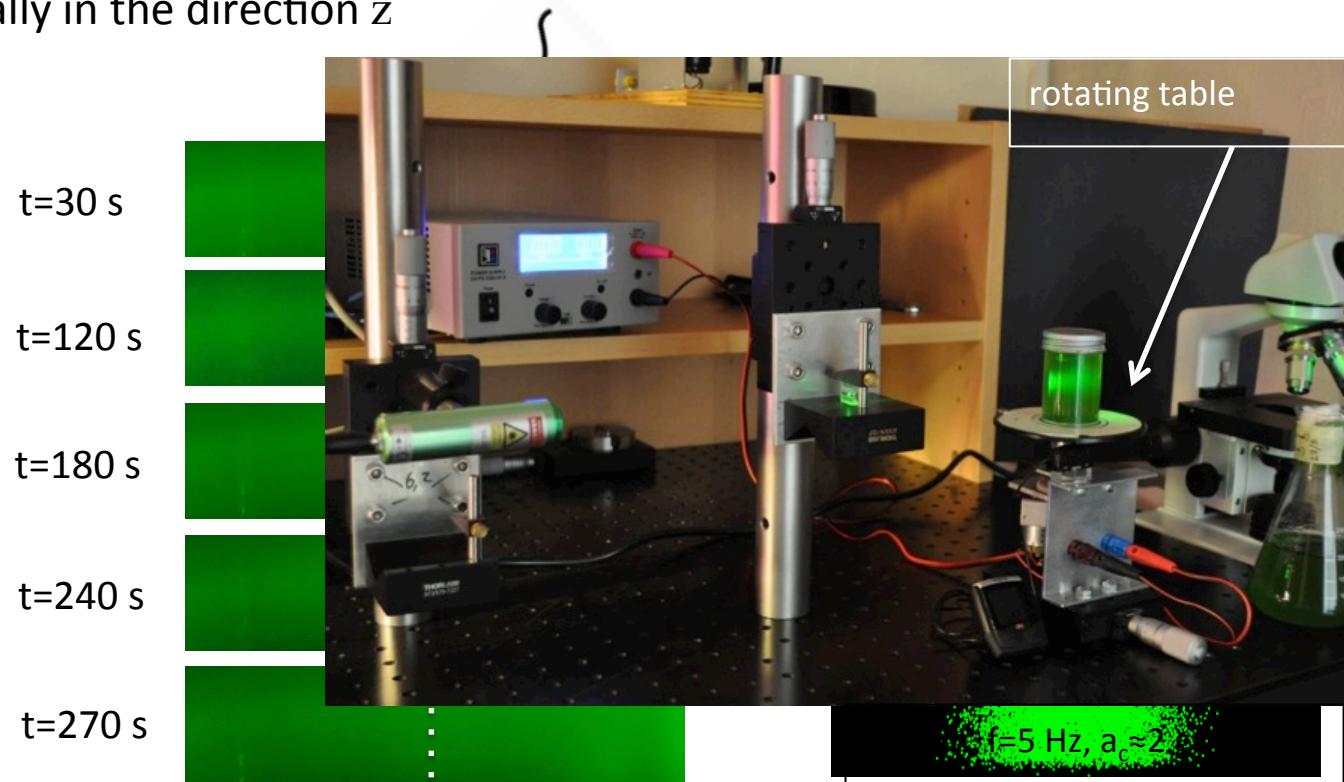
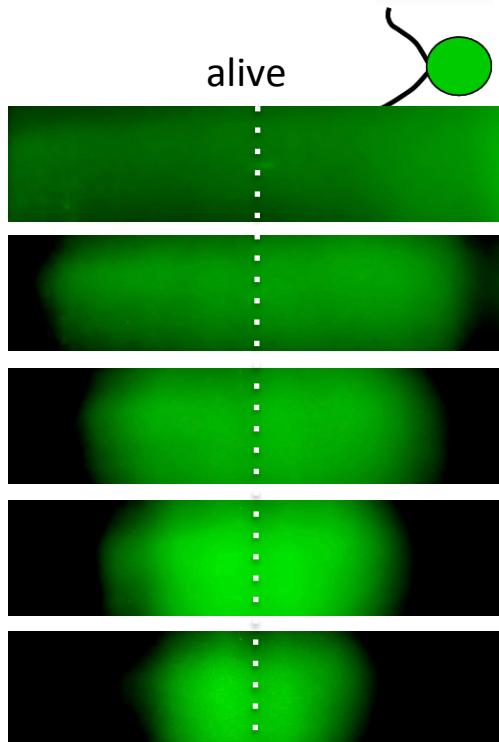
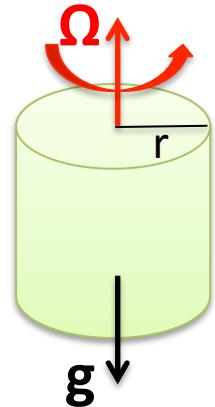
Cylinder in solid body rotation

$$\mathbf{u}(\mathbf{x}) = (-\Omega y, \Omega x, 0) \quad \omega = (0, 0, 2\Omega) \quad \mathbf{a}(\mathbf{x}) = (-\Omega^2 x, -\Omega^2 y, 0)$$

equilibrium swimming direction (stationary)  $\frac{d\mathbf{p}}{dt} = 0$

cylindrical coordinates  $\begin{cases} \mathbf{p}_r = -\gamma \mathbf{r} \\ p_z = \sqrt{1 - \gamma^2 r^2} \end{cases} \quad \gamma = \Omega^2/g$

gives the trajectory  $r(t) = r(0)e^{-\gamma v_s t}$  cells concentrate around the axis  
and  $\mathbf{p}$  aligns asymptotically in the direction z



# Simulations of turbulence

Simulation of the complete set of equations

Three dimensionless numbers

$$Re_\lambda = \frac{u_{rms} \lambda}{\nu} \quad \text{controls the weight of turbulent acceleration}$$

$$\Phi = \frac{v_s}{v_k} \quad \text{swimming number}$$

$$\Psi = \frac{\omega_{rms} v_o}{A_{rms}} \quad \text{stability number}$$

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} = -\nabla p + \nu \nabla^2 \mathbf{u} + \mathbf{f}$$

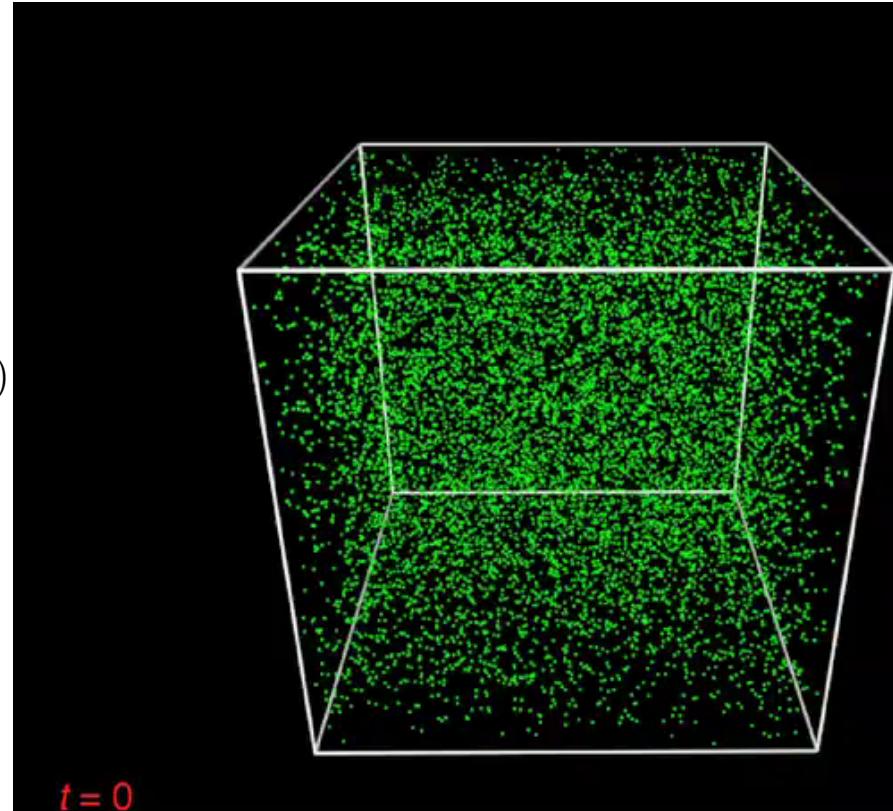
$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + v_s \mathbf{p}$$

$$\frac{d\mathbf{p}}{dt} = \frac{1}{2v_o} [\mathbf{A} - (\mathbf{A} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$

Gyrotactic swimmers as a dissipative system

$$\sum_{i=1}^d \left( \frac{\partial \dot{X}_i}{\partial X_i} + \frac{\partial \dot{\mathbf{p}}_i}{\partial \mathbf{p}_i} \right) = -\frac{d-1}{2v_o} (g\mathbf{p}_z + \mathbf{a} \cdot \mathbf{p})$$

as  $\mathbf{p}$  orients in the direction of local acceleration, swimming cells concentrate on a (fractal) subset of the phase space



# Turbulence at small Reynolds numbers ( $a_{rms} \ll g$ )

Typical conditions in the ocean mixing layer  $\varepsilon = 10^{-7} m^2 s^{-3}$

$$\eta = (\nu^3 / \varepsilon)^{1/4} \simeq 2 mm$$

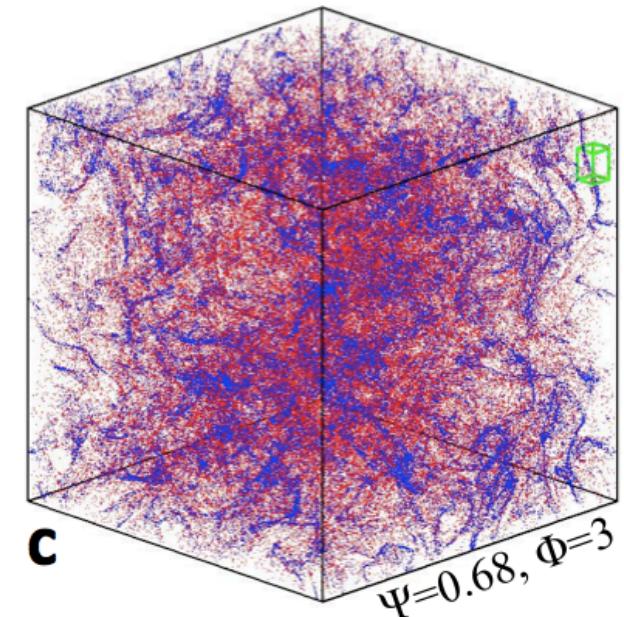
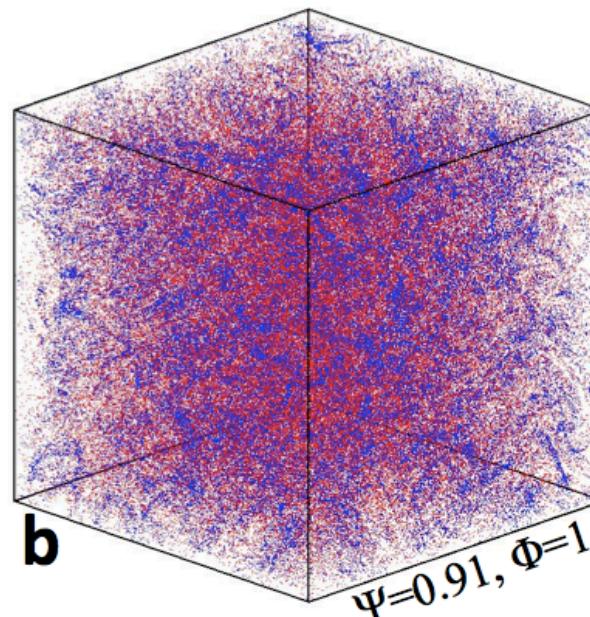
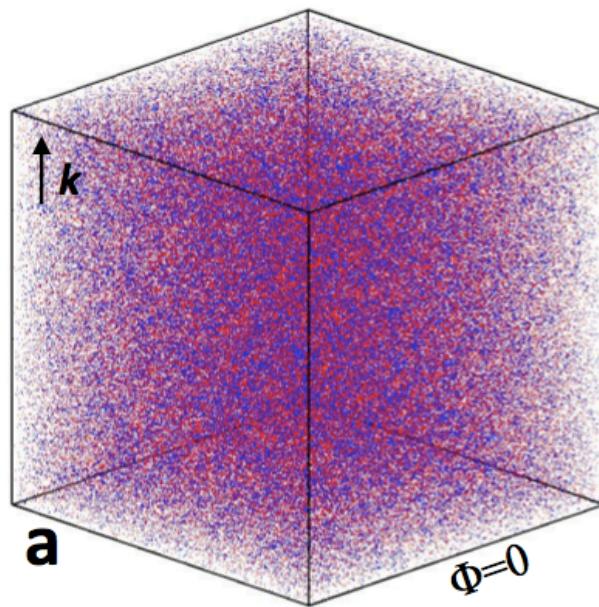
$$\tau_k = (\nu / \varepsilon)^{1/2} \simeq 3 s$$

$$v_k = (\nu \varepsilon)^{1/4} \simeq 0.5 mm s^{-1}$$

$$\Phi = \frac{v_s}{v_k} \simeq 0.4$$

$$\Psi = \frac{\omega_{rms} v_o}{g} \simeq 0.3$$

$$Re_\lambda = 65$$



$$n > 2\langle n \rangle$$

How clustering depends on parameters ?  
Where do cells cluster ?

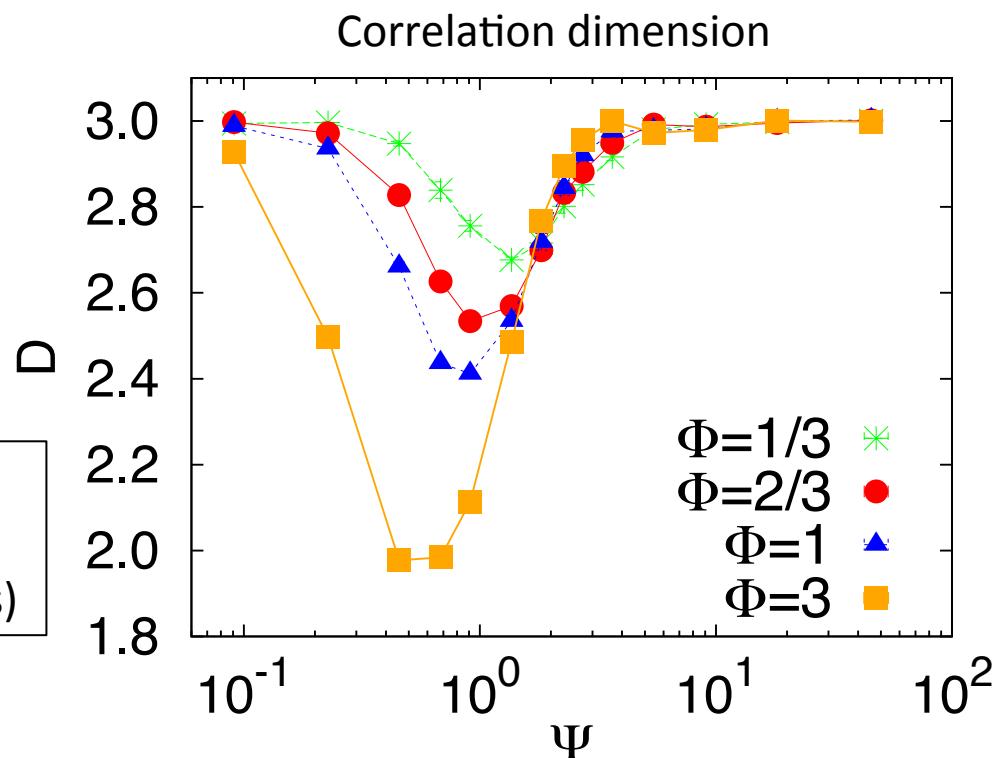
$10^6$  cells

# Fractal clustering

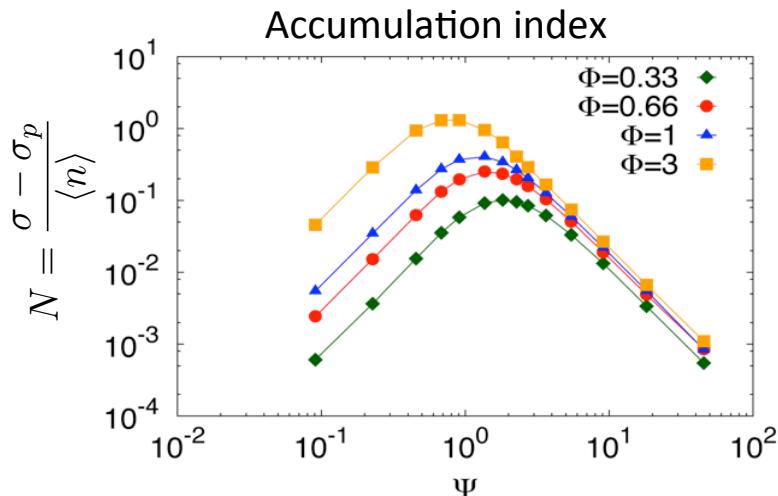
$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + \Phi \mathbf{p}$$

$$\frac{d\mathbf{p}}{dt} = \frac{1}{2\Psi} [\mathbf{k} - (\mathbf{k} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$

Homogeneous distribution in both limits  $\Psi \rightarrow 0$  (uniform vertical swimming) and  $\Psi \rightarrow \infty$  (swimming in random directions)



Clustering is maximum (D minimum) for  $\Psi \simeq 1$  and increases with  $\Phi$



$$N = \frac{\sigma - \sigma_p}{\langle n \rangle}$$

$$\sigma = (\langle n^2 \rangle - \langle n \rangle^2)^{1/2}$$

swimming number	$\Phi = \frac{v_s}{v_k}$
stability number	$\Psi = \frac{\omega_{rms} v_o}{g}$

# How clustering depends on $\Psi$

$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + \Phi \mathbf{p}$$

$$\frac{d\mathbf{p}}{dt} = \frac{1}{2\Psi} [\mathbf{k} - (\mathbf{k} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$

For small  $\Psi$ , at first order we have

$$\mathbf{p} = (\Psi\omega_y, -\Psi\omega_x, 1)$$

passive tracers in an effective velocity field

$$\mathbf{v} = \mathbf{u} + \Phi \mathbf{p}$$

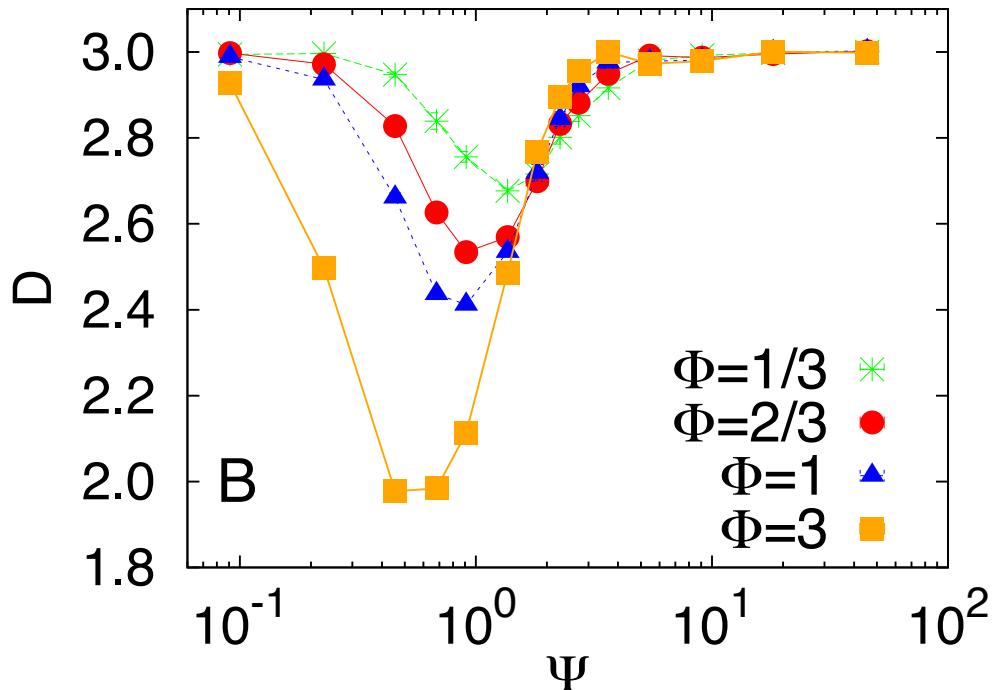
with divergence

$$\nabla \cdot \mathbf{v} = \Phi \nabla \cdot \mathbf{p} = -\Psi \Phi \nabla^2 u_z$$

Fractal codimension  $3 - D \propto (\nabla \cdot \mathbf{v})^2$

and therefore

$$D = 3 - a(\Phi\Psi)^2$$



A weakly compressible velocity field

$$\mathbf{v} = \mathbf{u} + \delta \mathbf{w}$$

$$\text{with } \nabla \cdot \mathbf{u} = 0 \quad \nabla \cdot \mathbf{w} \neq 0$$

tracers cluster on a fractal set with codimension given by  $d - D_2 \simeq \delta^2$

G Falkovich, A Fouxon, MG Stepanov,  
*Nature* **419**, 151-154 (2002).

I Fouxon, *Phys. Rev. Lett.* **108**, 134502 (2012).

## How clustering depends on $\Psi$

$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + \Phi \mathbf{p}$$

$$\frac{d\mathbf{p}}{dt} = \frac{1}{2\Psi} [\mathbf{k} - (\mathbf{k} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$

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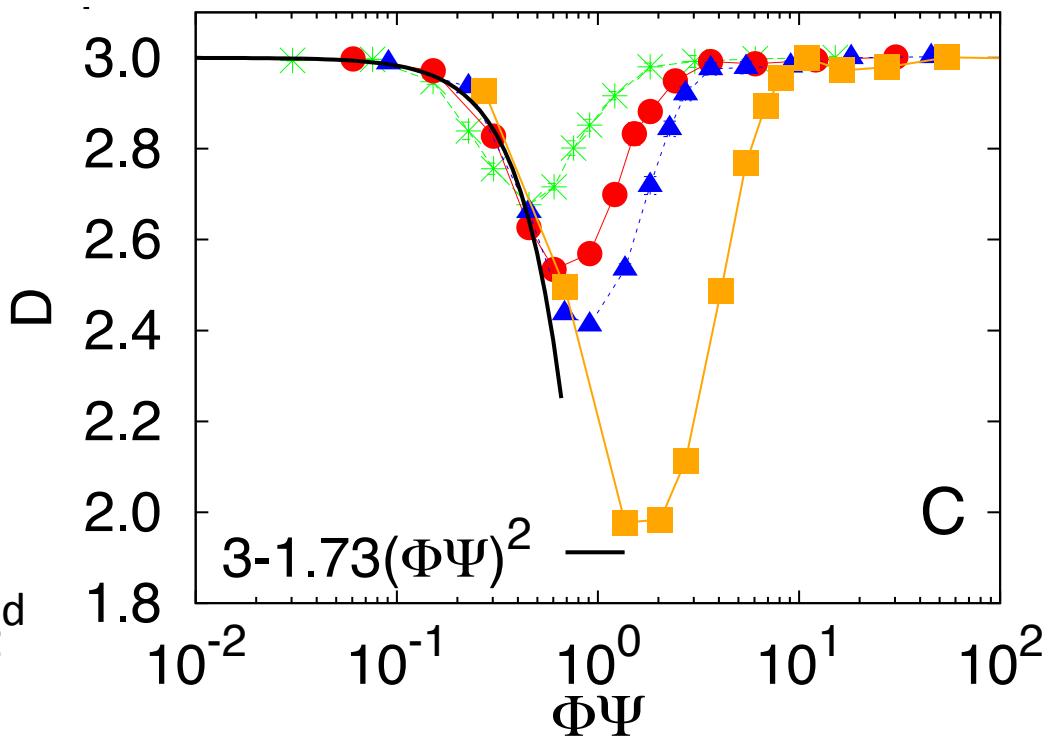
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# Where do cells cluster ?

Swimmers as tracers transported by a weakly compressible flow

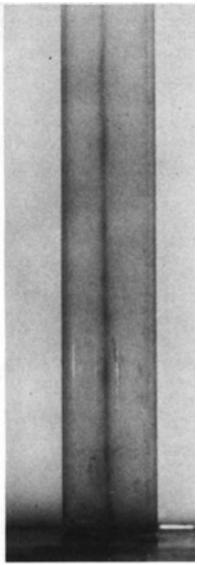
$$\nabla \cdot \mathbf{v} = \Phi \nabla \cdot \mathbf{p} = -\Psi \Phi \nabla^2 u_z$$

concentrate on regions where  $\nabla^2 u_z > 0$

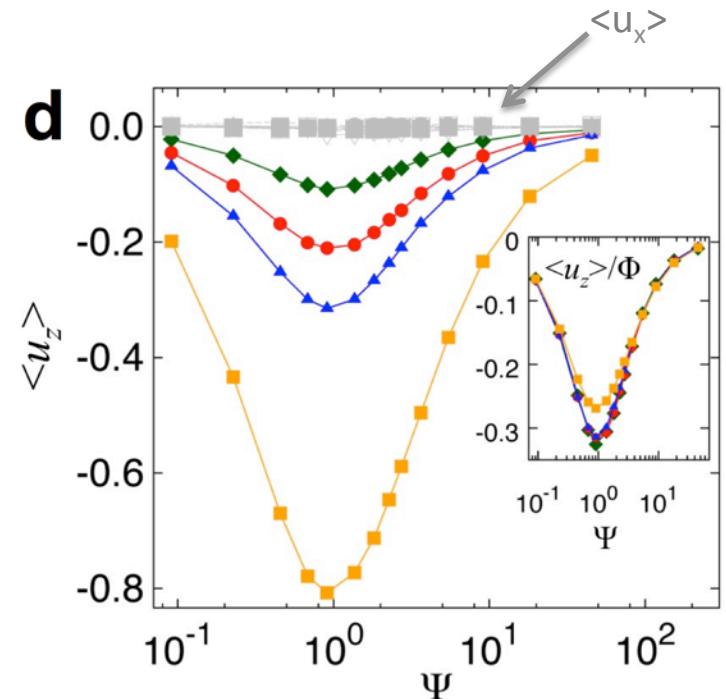
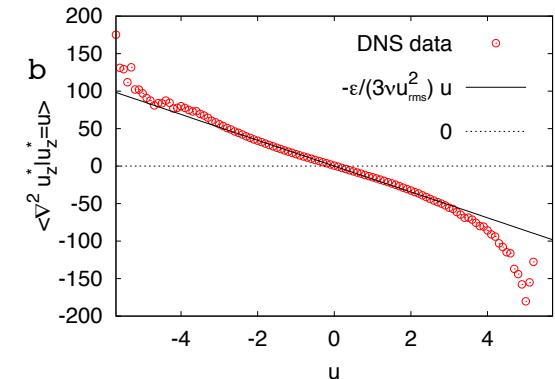
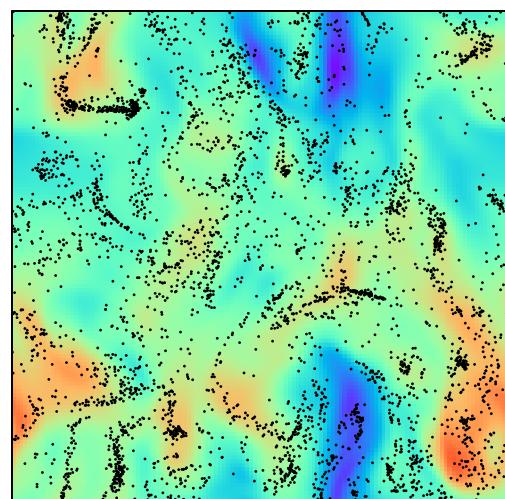
In homogeneous, isotropic turbulence

$$\epsilon = \nu \langle (\nabla \mathbf{u})^2 \rangle = -3\nu \langle u_z \nabla^2 u_z \rangle$$

and therefore  $\nabla^2 u_z > 0$  means  $u_z < 0$



Swimming cells  
accumulate in  
downwelling  
regions,  
where  $u_z < 0$



# Clustering at increasing Reynolds numbers

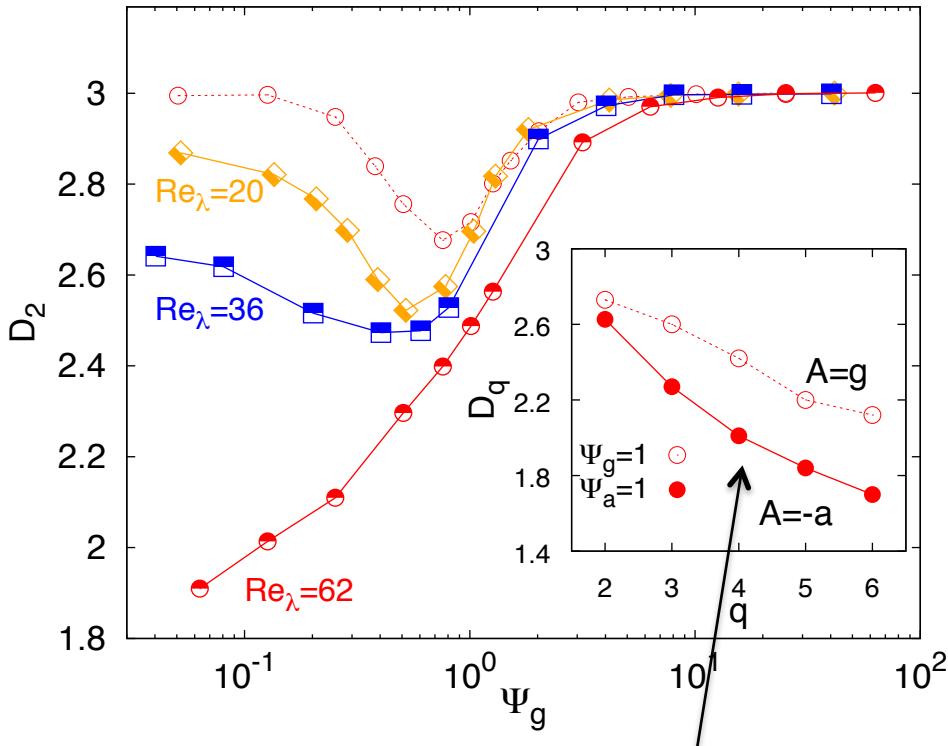
$\Phi = 1/3$

- Clustering increases with Reynolds
- Minimum of D disappears for large Re  
(when  $\alpha \approx 1$ )

$$\alpha \equiv \frac{a_{rms}}{g}$$

$$\begin{aligned} \alpha &= 0.34 \\ \alpha &= 0.50 \\ \alpha &= 0.84 \end{aligned}$$

$$\begin{aligned} Re_\lambda &= 20 \\ Re_\lambda &= 36 \\ Re_\lambda &= 62 \end{aligned}$$



Turbulent accelerations enhance cell clustering

multifractal clustering

At small  $\Psi$  the swimming direction aligns towards strong acceleration regions which are not uniform

What is the role of fluid acceleration ?

# Clustering in the limit $g=0$

To understand the role of acceleration we consider the case  $a_{\text{rms}} \gg g$  and take  $g=0$

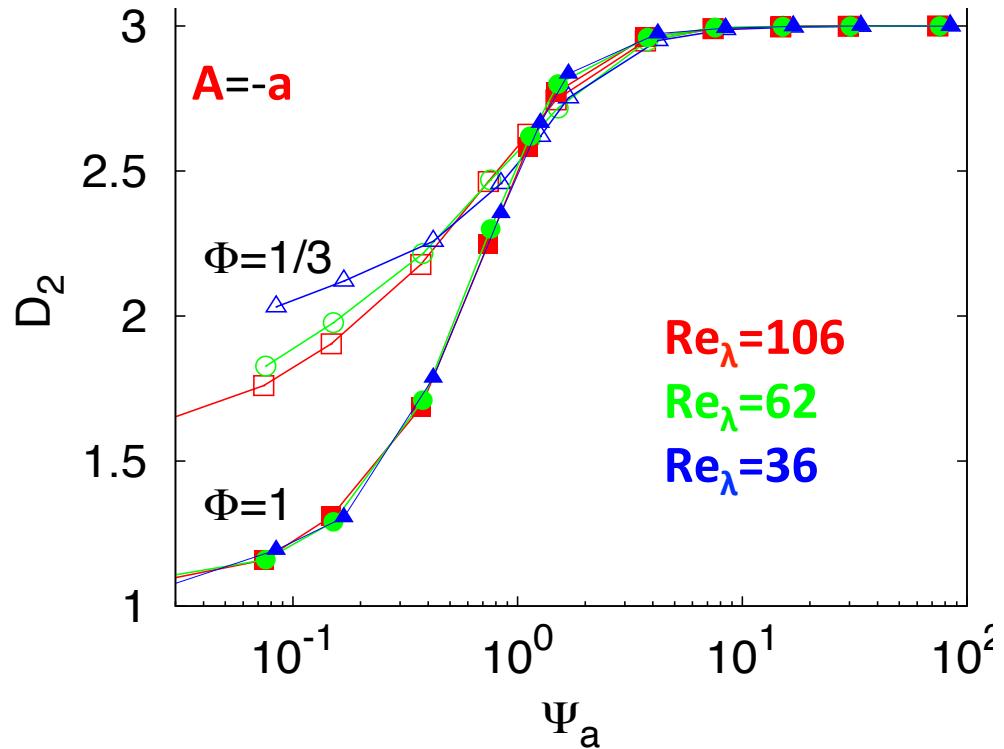
$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + \Phi \mathbf{p}$$

$$\frac{d\mathbf{p}}{dt} = \frac{1}{2\Psi} [\mathbf{a} - (\mathbf{a} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$

clustering increases with cell stability  
i.e when  $\Psi \ll 1$

clustering increases with swimming speed i.e with  $\Phi$

weak (if any) dependence on  $Re_\lambda$



**what is driving clustering? where do cells go?**

# Predictions for small $\Psi$

$\Psi \ll 1$ ,  $\mathbf{p}$  is mainly aligned with  $\mathbf{a}$ .

Effective velocity for swimmers  $\mathbf{v} \simeq \mathbf{u} + \Phi \hat{\mathbf{a}}$

a compressible field with

$$\nabla \cdot \mathbf{v} \simeq \Phi \nabla \cdot \hat{\mathbf{a}}$$

Clearly  $\nabla \cdot \hat{\mathbf{a}} \not\propto \nabla \cdot \mathbf{a}$  however their **sign** is strongly **correlated**. Swimmers accumulate where

$$\nabla \cdot \mathbf{a} < 0$$

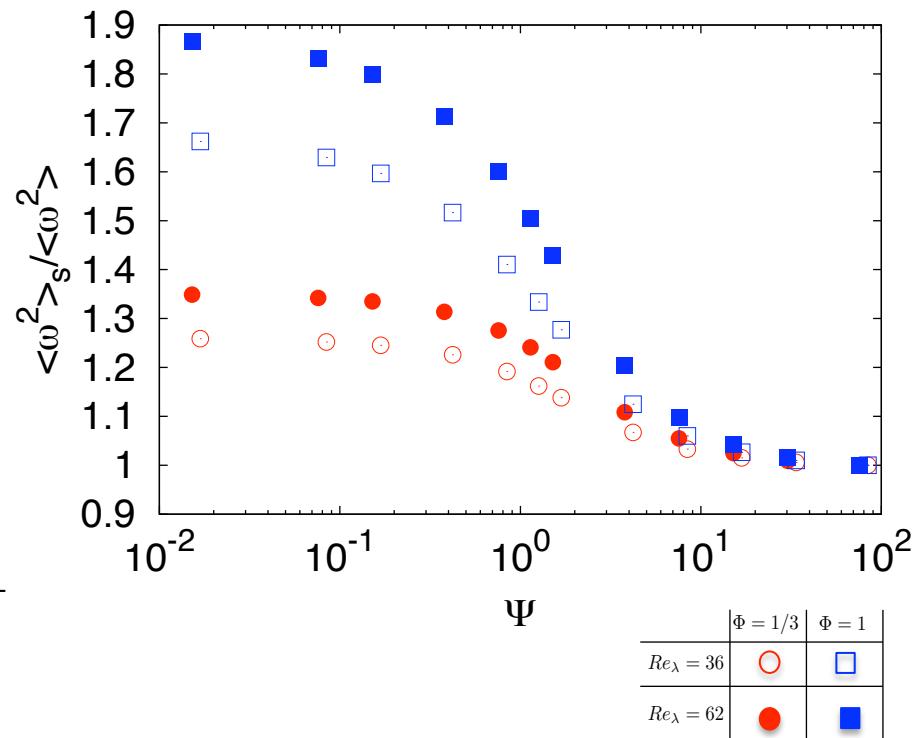
$$\nabla \cdot \mathbf{a} = \sum_{ij} (\hat{S}_{ij}^2 - \hat{\Omega}_{ij}^2)$$

$$\hat{S}_{ij} = \frac{\partial_j u_i + \partial_i u_j}{2} \quad \hat{\Omega}_{ij} = \frac{\partial_j u_i - \partial_i u_j}{2}$$

$\nabla \cdot \mathbf{a} < 0$  corresponds to large vorticity regions

$$\frac{d\mathbf{x}}{dt} = \mathbf{u}(\mathbf{x}, t) + \Phi \mathbf{p}$$

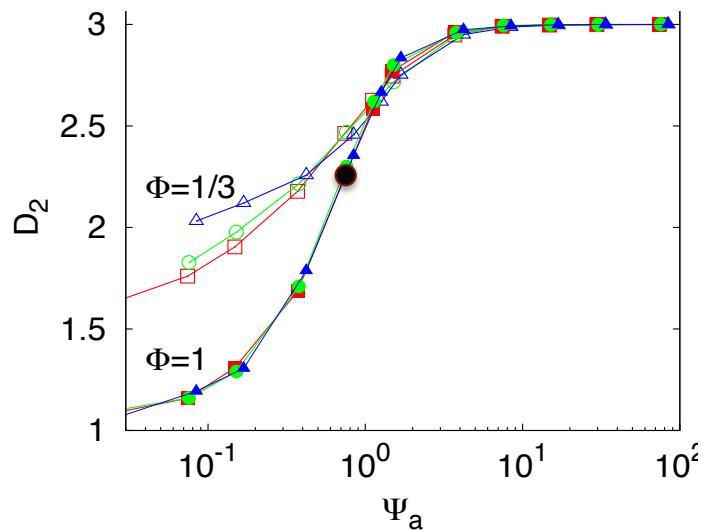
$$\frac{d\mathbf{p}}{dt} = \frac{1}{2\Psi} [\mathbf{a} - (\mathbf{a} \cdot \mathbf{p})\mathbf{p}] + \frac{1}{2}\boldsymbol{\omega} \times \mathbf{p}$$



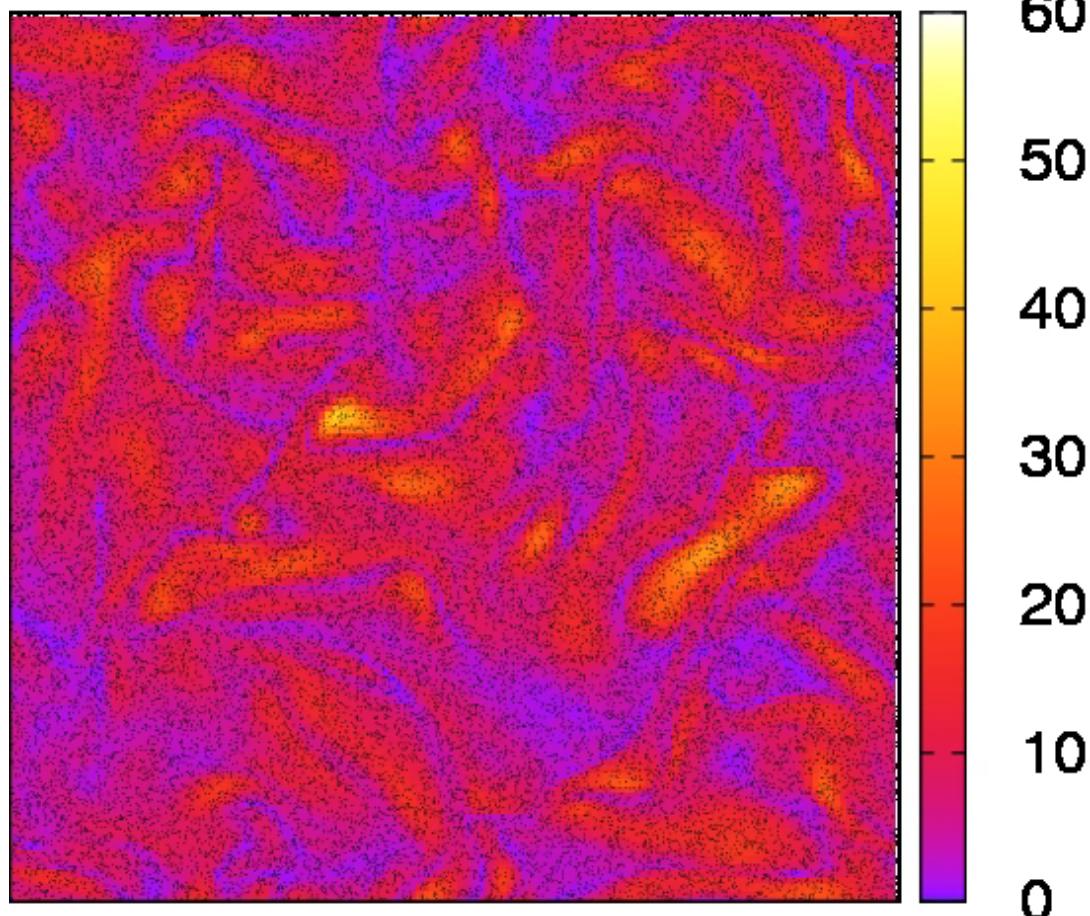
**Swimmers concentrate in vortices** (like light particles)

$\langle \dots \rangle_s$  average at swimmer positions

# Swimmers concentrate in vortices



$$\begin{aligned}Re_\lambda &= 62 \\ \Psi &= 1.5 \\ \Phi &= 1\end{aligned}$$



# A prediction for $D_2$

Clustering is more efficient for faster swimmers

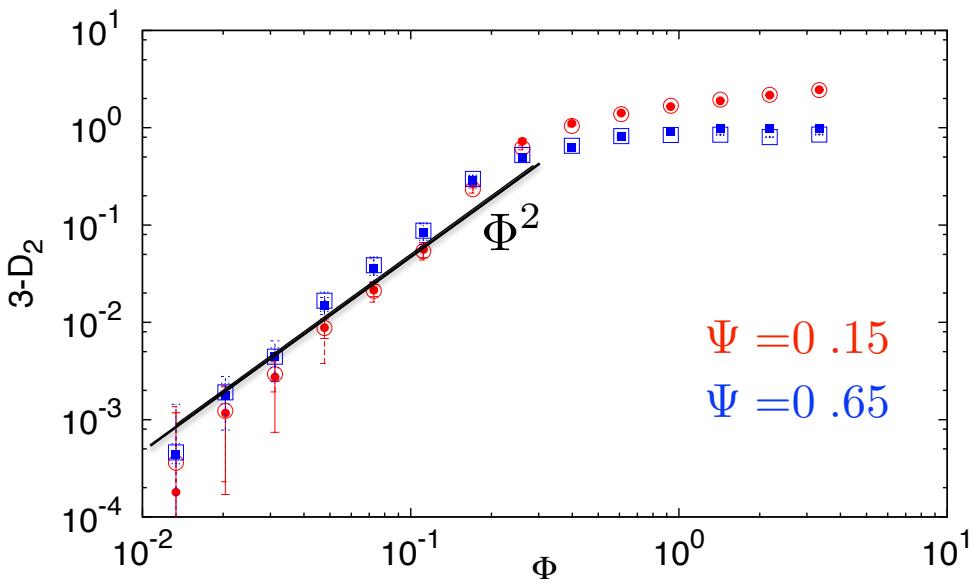
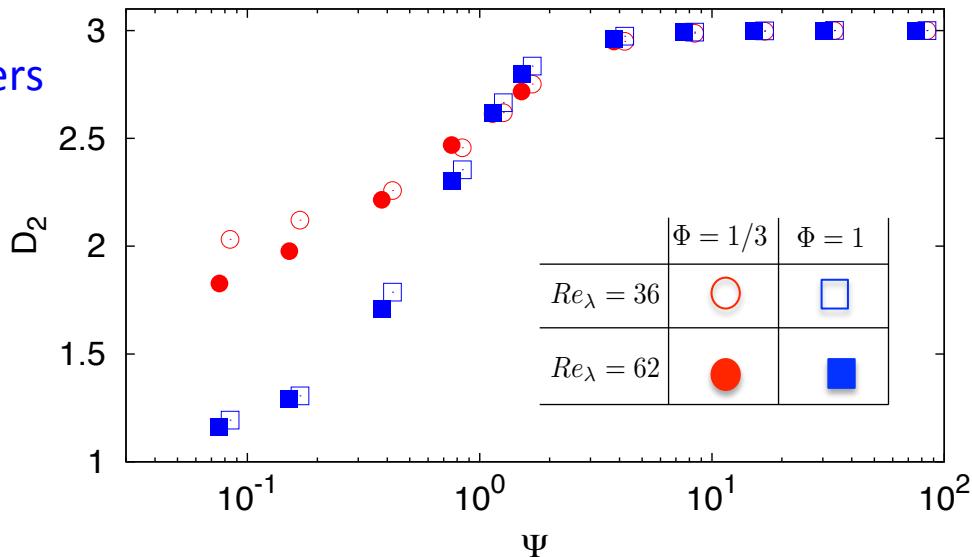
How does  $D_2$  depends on  $\Phi$  ?

Effective velocity for swimmers

$$\mathbf{v} \simeq \mathbf{u} + \Phi \hat{\mathbf{a}}$$

$$\text{with } \nabla \cdot \mathbf{v} \propto \Phi \nabla \cdot \hat{\mathbf{a}}$$

$$3 - D_2 \simeq \Phi^2$$



Happy  
Birthday  
Angelo